Characterization and Modeling of Naturally Fractured Reservoirs: The Role of Outcrop Analogues

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Acknowledgements to the Stanford Rock Physics Laboratory (SRB), Stanford Rock Fracture Project (RFP), and Repsol-YPF (Bolivia)



Characterization and Modeling of Fractured Reservoirs: the Big Picture

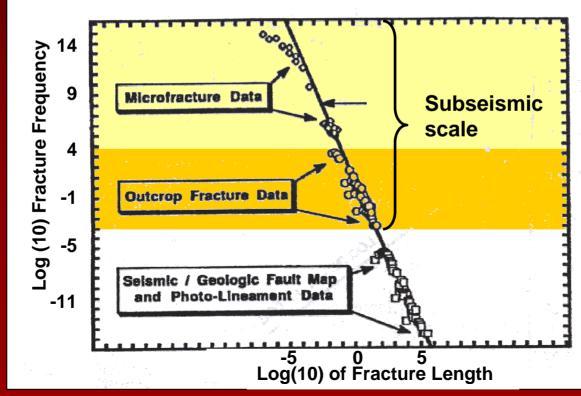
iReservoir's Integrated Approach

- 1. Create Equally Probable Geologic Models
 - *Quantify* spatial heterogeneity
 - Identify the *fracture indicators* (curvature, faults, strain, lithology, etc.)
 - Map crack density from fracture indicators
 - Create site-specific equally probable scenarios
- 2. Constrain Geologic Models using *Seismic, Log and Engineering Data*
- 3. Use *History Matching* to constrain and validate the geologic model



Why Outcrop Descriptions?

- 1. Spatial heterogeneity at subseismic scale
- 2. Deformation mechanisms and parameters controlling fracture distribution
- 3. Reservoir specific statistics





After Heffer & Bevan (1990)

Outline

- 1. Why outcrop analogues ?
- 2. Deformation mechanisms from outcrops
- 3. Example of faults and fractures systems fold and thrust belts:
 - Hierarchical shearing and progressive deformation,
 - Fracture hierarchies and stratigraphic architecture,
 - Factors controlling fracture distribution.
 - Comparison to subsurface.
- 4. The role of conjugate faulting on fracture localization.
- 5. Use of outcrop data for geostatistical modeling.
- 6. Conclusions.



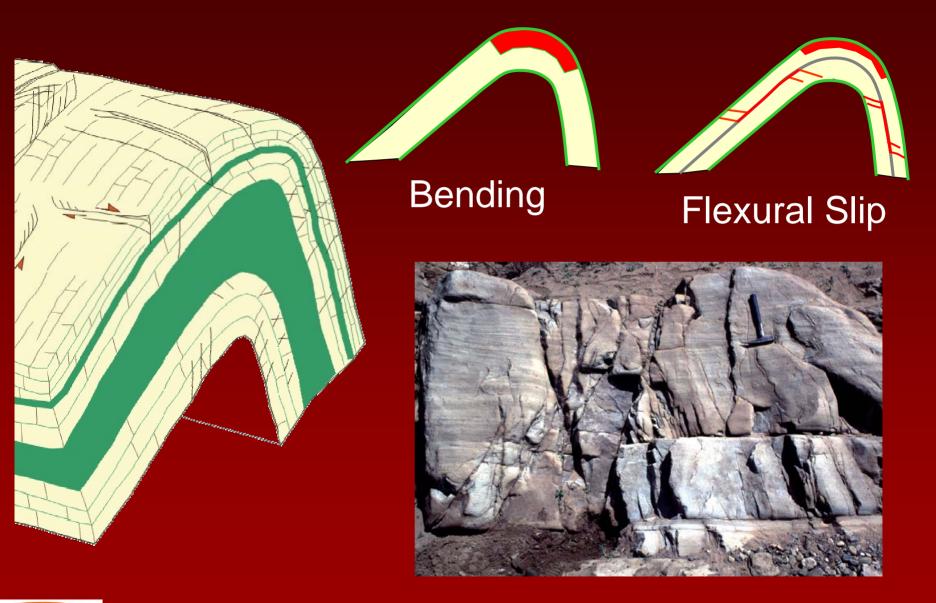
Deformation Mechanisms from Outcrops





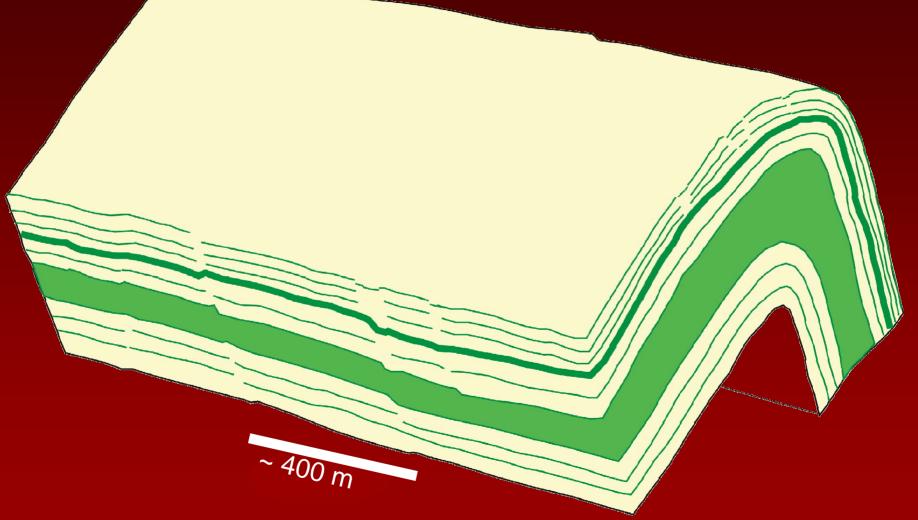
Florez, 2006

Fracturing Mechanisms





Faults and Fractures in a Fold and Thrust Belt

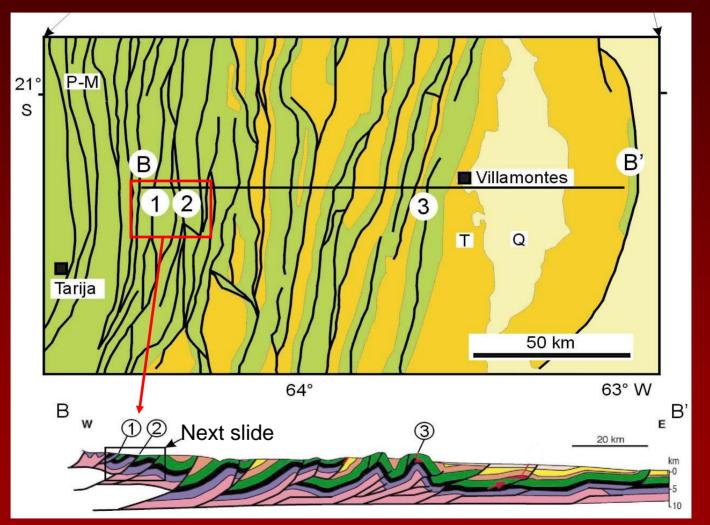


How are faults and fractures distributed in this Anticline?



Florez, 2006

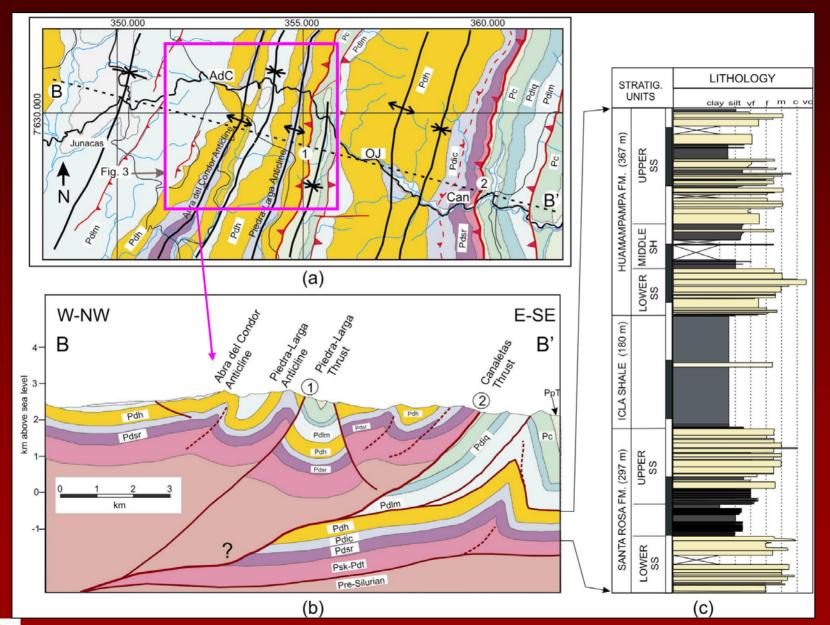
Bolivian Subandean Fold and Thrust Belt



After Labaume et al, 2001



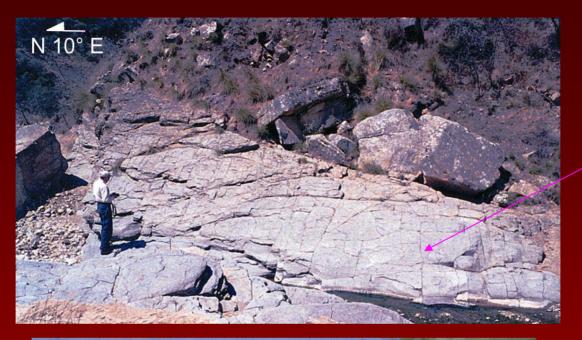
Faults and Fractures at the Abra del Condor Anticline



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Joints and Sheared Joints





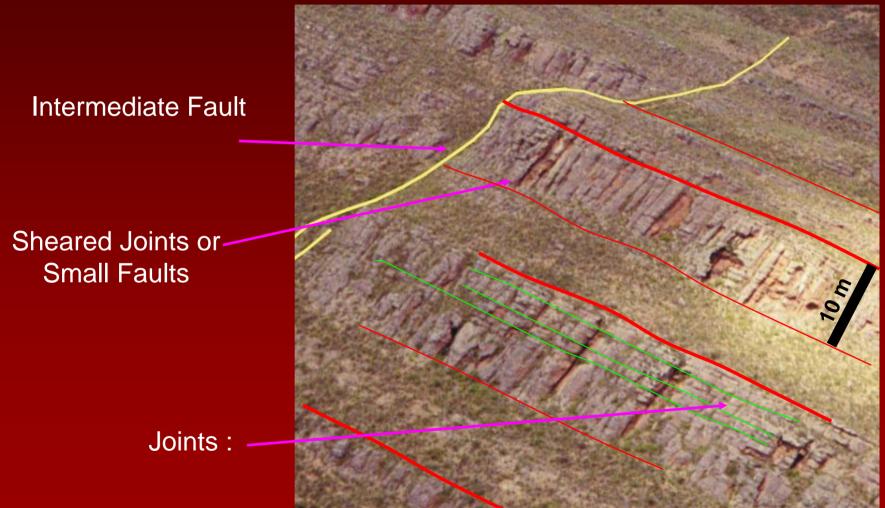


Sheared Joints



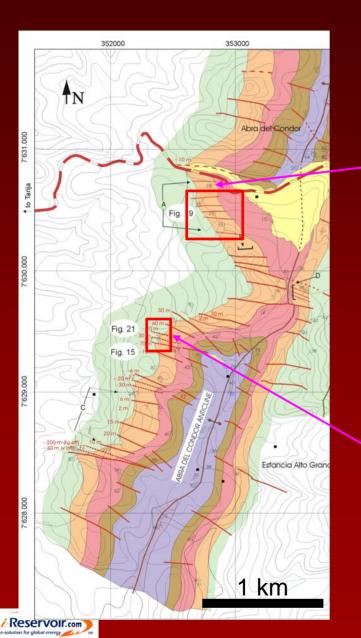
Small and Intermediate Faults

Faults and fractures in this case are the result of hierarchical shearing and progressive deformation. Four main hierarchies identified: (1) joints, (2) small faults, (3) intermediate faults, (4) fault zones.

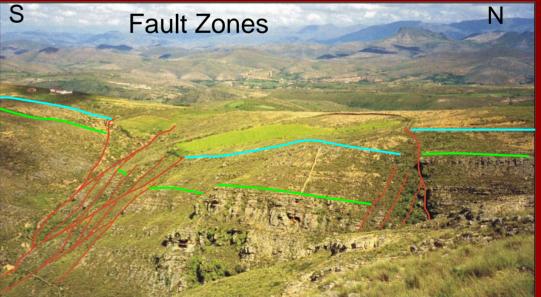




Fault Zones and Intermediate Faults

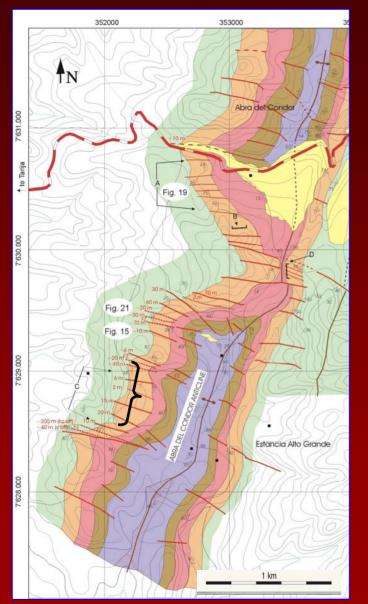


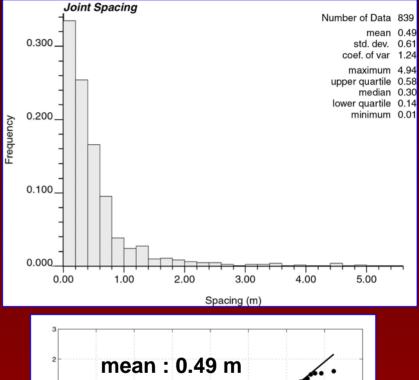


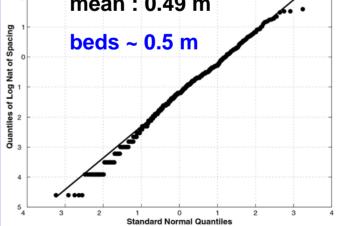


Florez, 2006

Spacing of Joints

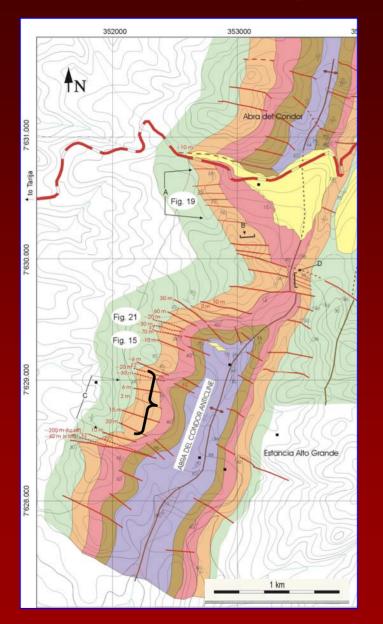


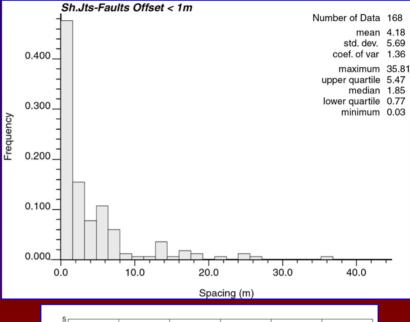


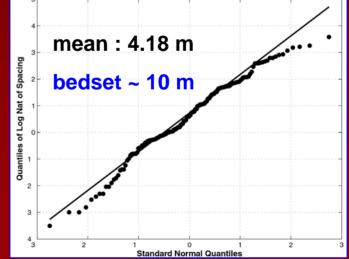




Spacing of Small Faults

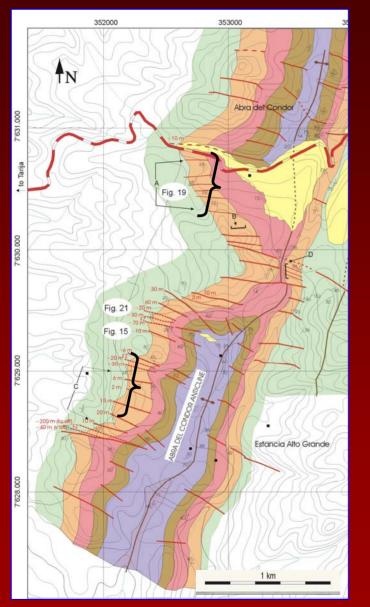


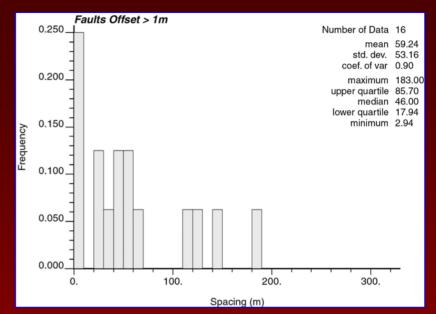


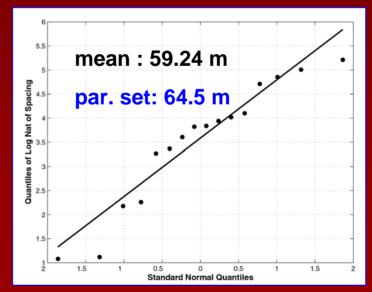




Spacing of Intermediate Faults

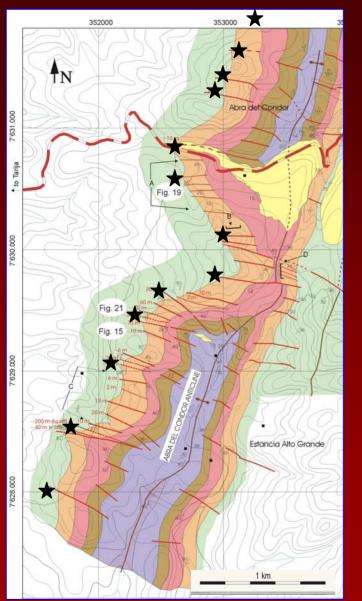


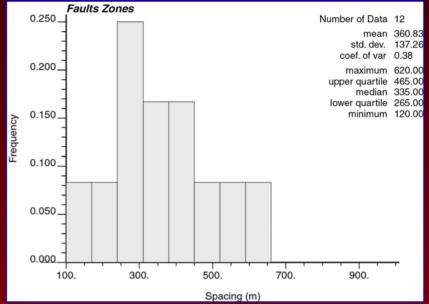


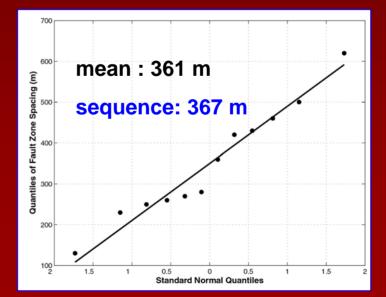


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Spacing of Fault Zones

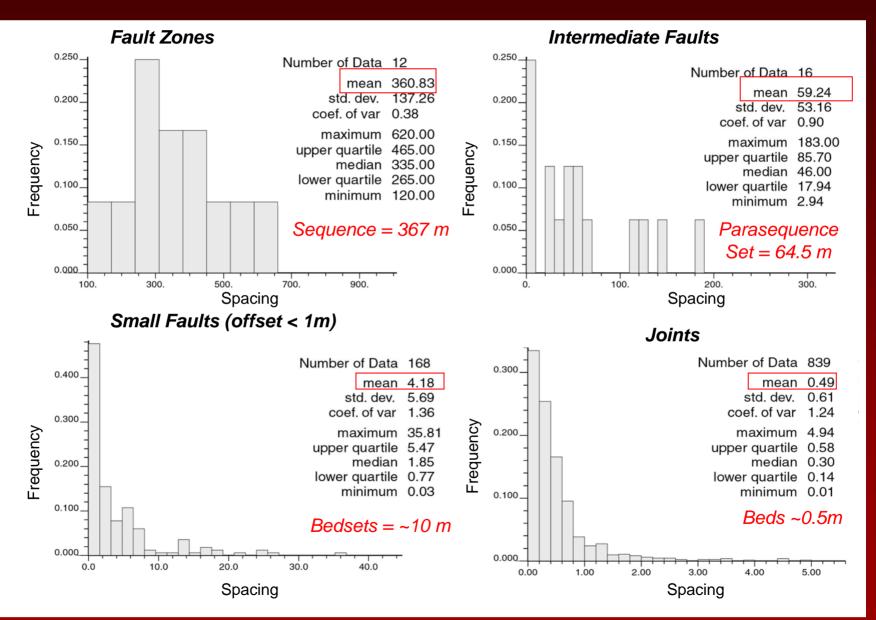






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Fracture Spacing and Stratigraphy



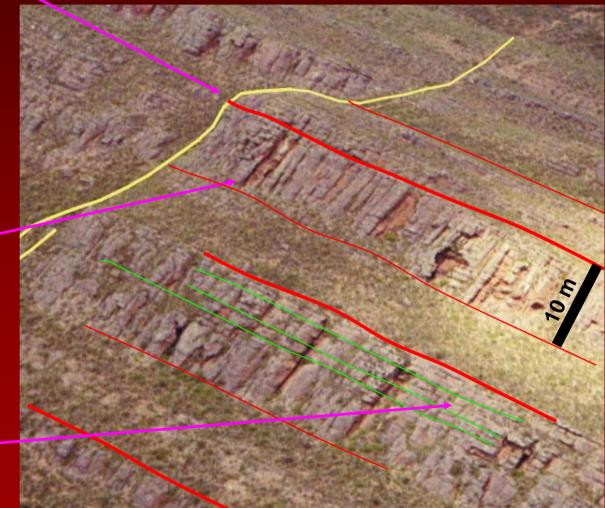


Factors Controlling Fracture Distribution: Stratigraphy and Lithofacies

Intermediate Fault: Parasequence set

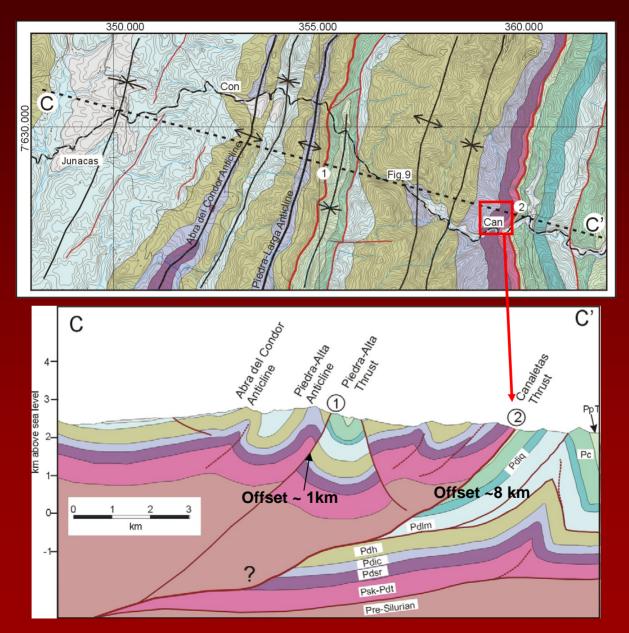
Sh-Joints/SF : Bedsets and Parasequences

> Joints : Beds





Fracture Spacing and Shear Strain





Fracture Spacing and Shear Strain

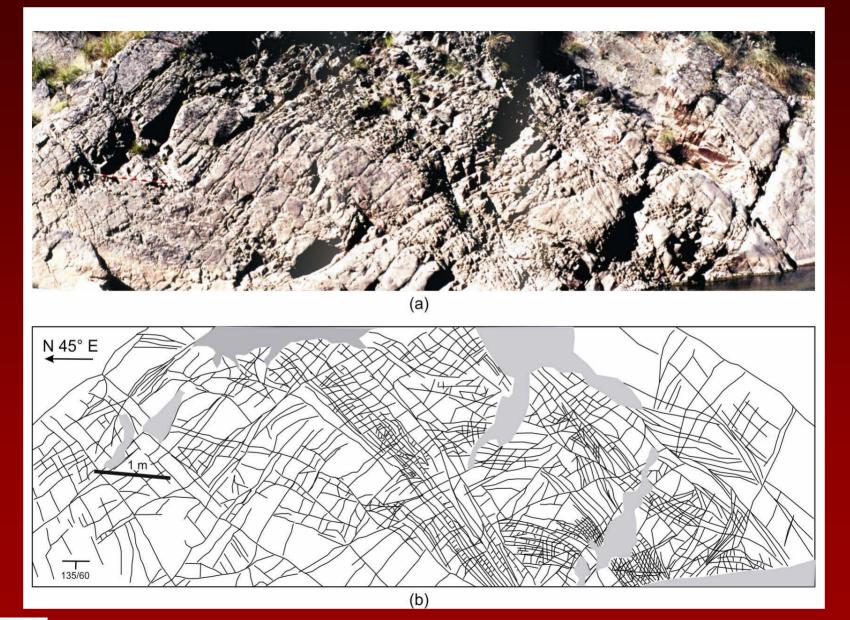
Small Faults; mean 0.76 m

Intermediate Faults; mean 19.3 m



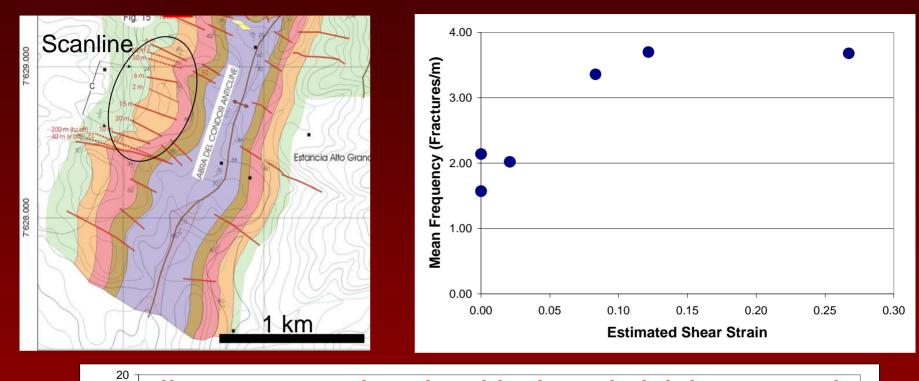


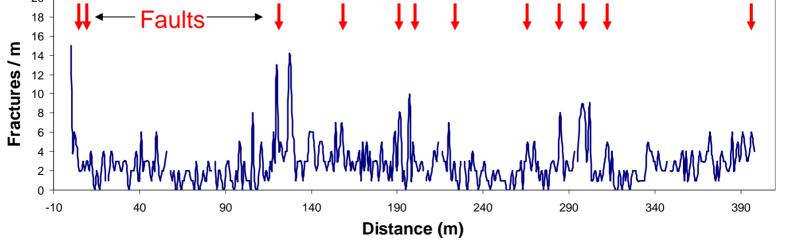
Complex Fracture Pattern near Canaletas Thrust





Shear Strain and Fracture Density

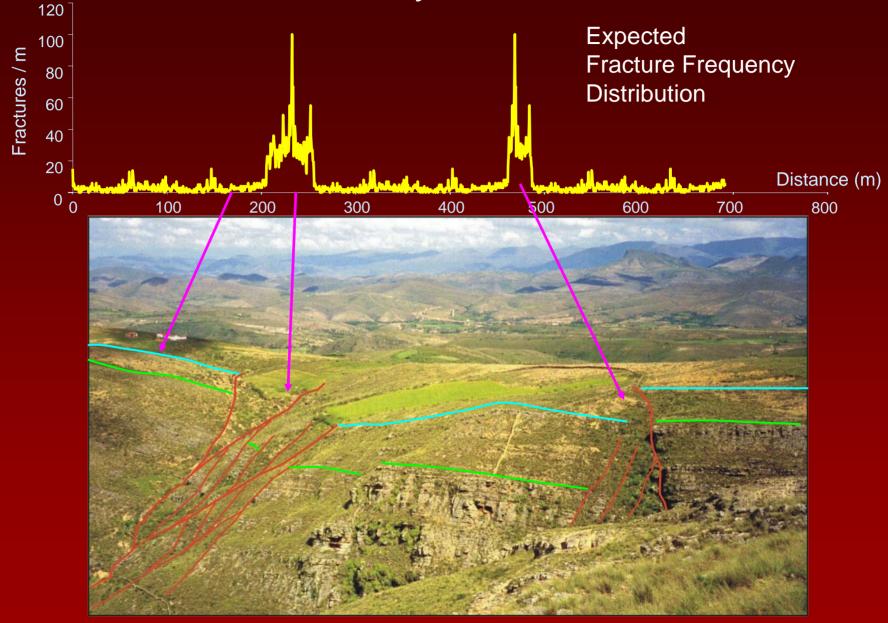






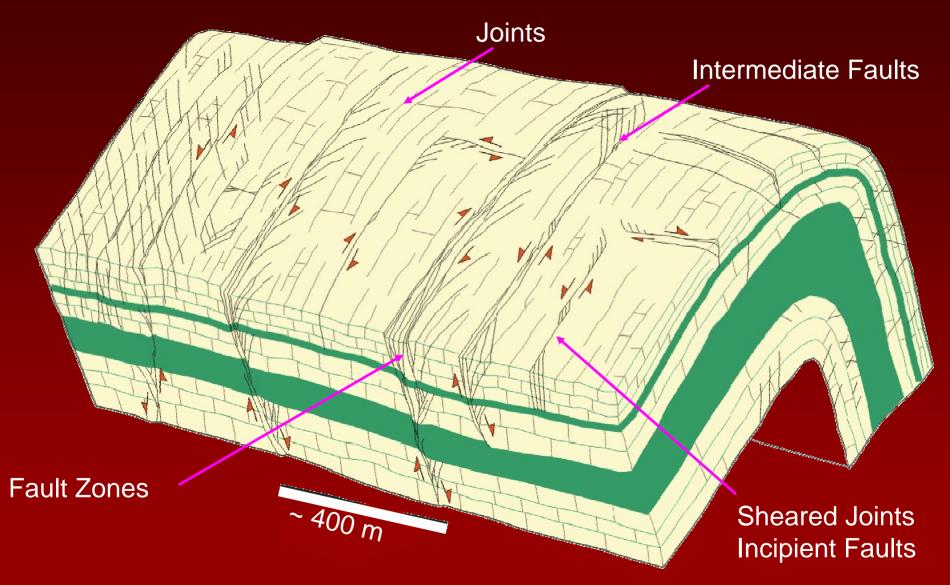
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Fracture Density and Fault Zones



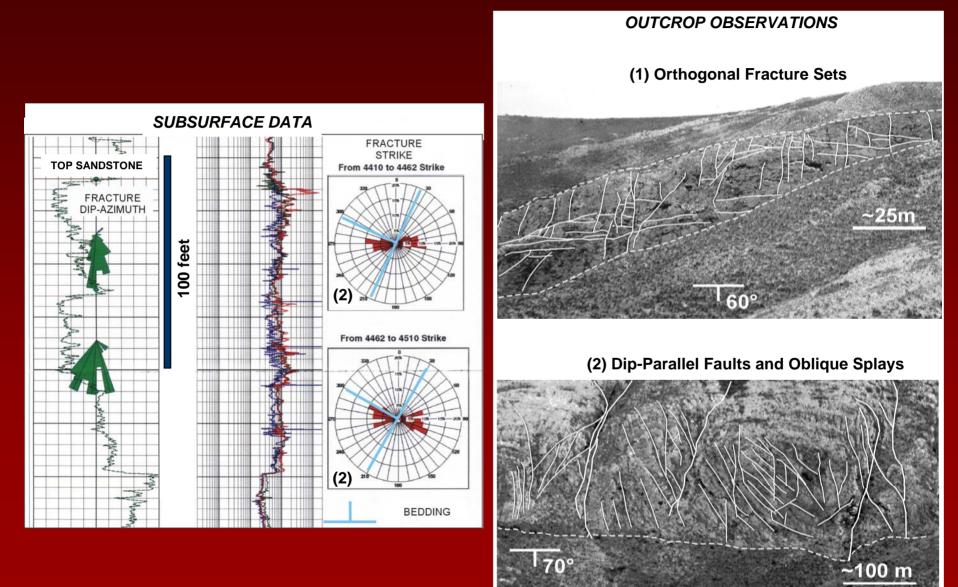


How are Faults are Fractures distributed?





Comparison to Subsurface: Orientation

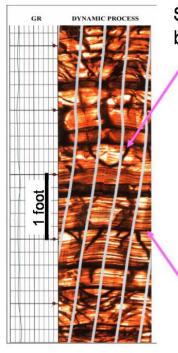


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Comparison of Fracturing Mechanisms

Flexural Slip

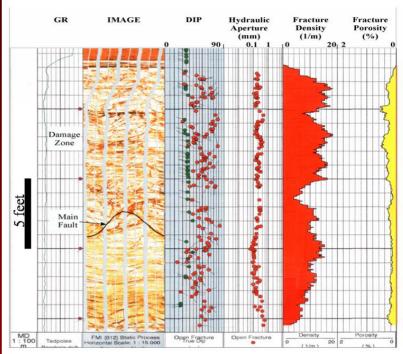


Splay joints abutting against bedding surfaces

Splay-confining

bedding surfaces

Fracture Localization —



Damage Zone



Main Fault



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Comparison with other Areas (Zagros, Iran)

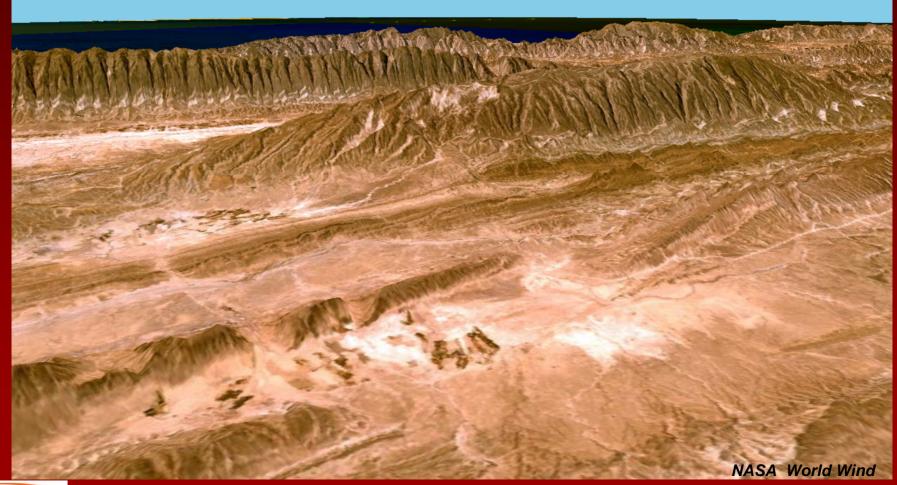




Florez, 2006

Oblique and Dip-Parallel Fracture Trends

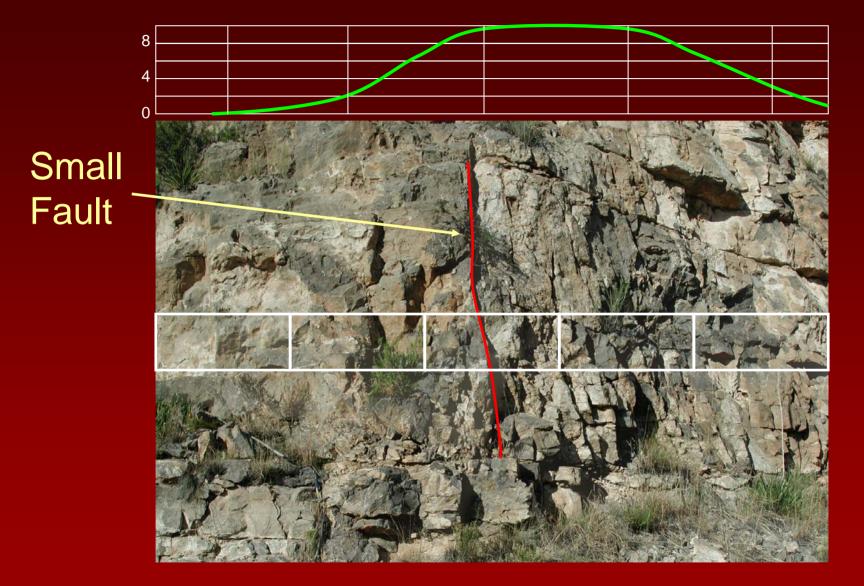
Latitude: 27.77408 Longitude: 52.88007 Heading: -123.65026 Tilt: 77.17896 Altitude: 7729m Distance: 36580m FOV: 45.00000 Terrain Elevation: 631.00 meters





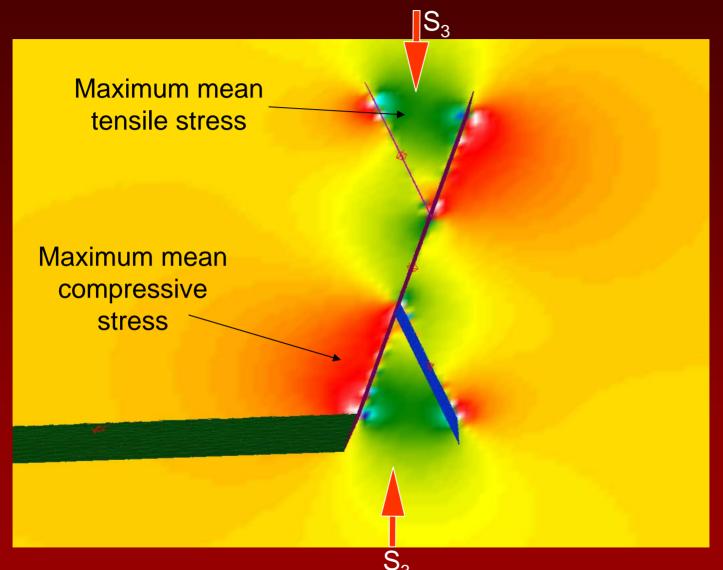
Florez, 2006

Conjugate Faults and Fracture Swarms





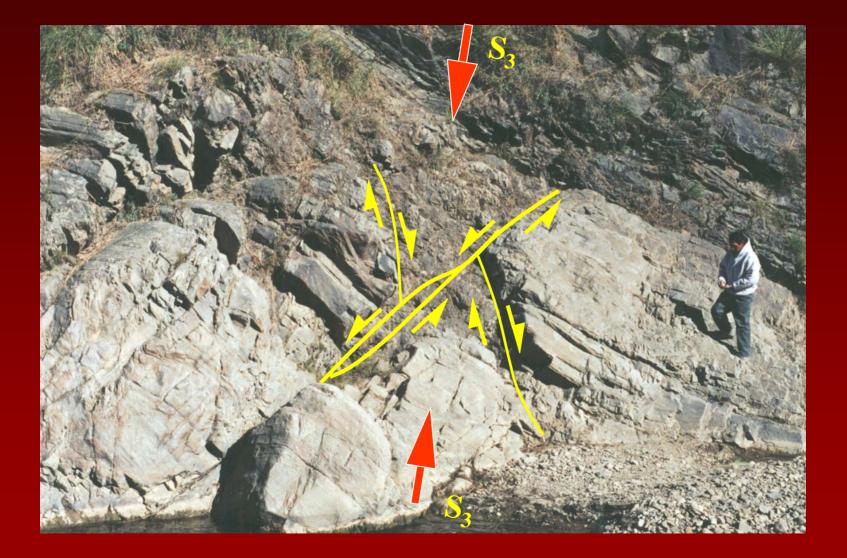
Conjugate Faults May Localize Tension



 S_3 Stress distribution due to slip along conjugate faults

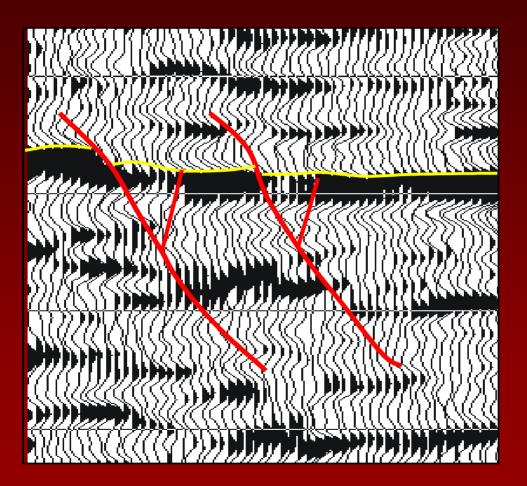


Fracture Swarm within Conjugate Faults





Implications for Seismic Interpretation



Asymmetric amplitude dimming around normal faults



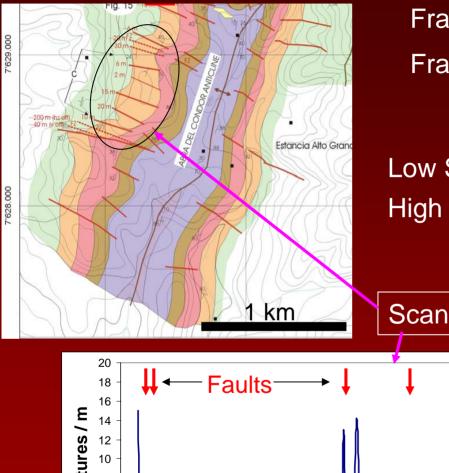
Florez, 2006

What we have learned so far:

- 1. Fractures occur at different scales as a result of hierarchical shearing and progressive deformation.
- 2. The spacing and dimension of different fracture hierarchies can be linked to stratigraphy
- 3. Shear strain and lithofacies are among the main factors controlling fracture density
- 4. Antithetic conjugate faults create fracture swarms



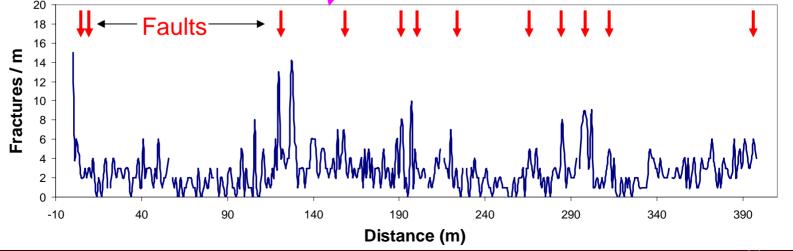
Fracture Modeling



Fracture Density = Fractures / m Fracture Density = 1 / Spacing

Low Shear-Strain = Low Fracture Density High Shear-Strain = High Fracture Density

Scanline at Low Shear-Strain Area





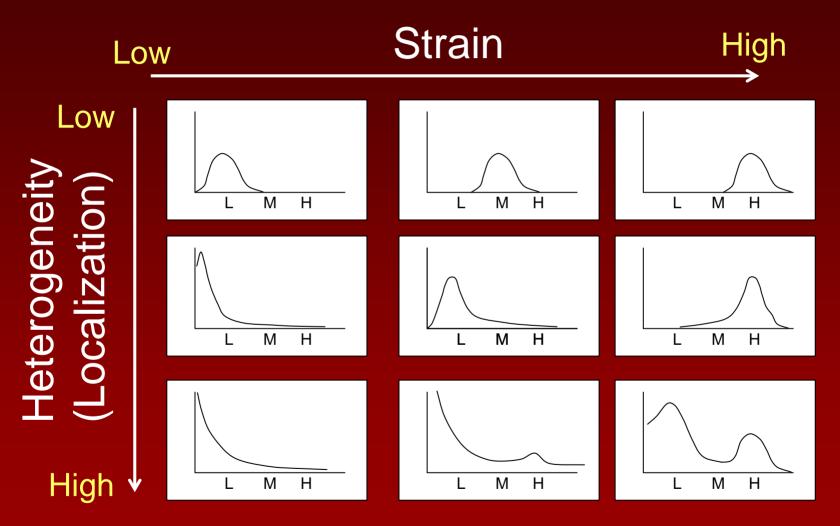
Fracture Density Distributions

Low Shear-Strain Area High Shear-Strain Area Normal Distribution Lognormal Distribution Number of Data 411 Number of Data 84 0.200 mean 3. 0.200 mean 29. std. dev. 2. std. dev. 11. coef. of var 1. coef. of var 0. 0.150 0.150 requency maximum 15. maximum 67. upper quartile 4. upper quartile 31. 0.100 median 2. 0.100 median 27. lower quartile 1. lower quartile 23. minimum 0. minimum 12. 0.050 0.050 0.000 0.000 0.0 10.0 20.0 30.0 100. 20. 40. 60 80. 0 Fractures/m Fractures/m **Combined Data Exponential Distribution** 0.400 Number of Data 495 mean 7. Low Shear-Strain Area 0.300 std. dev. 11. Frequency coef. of var 2. maximum 67. 0.200 upper quartile 5. median 3. High Shear-Strain Area lower quartile 2. 0.100 minimum 0. 0.000 20. 40. 60. 80. 100. Fractures/m



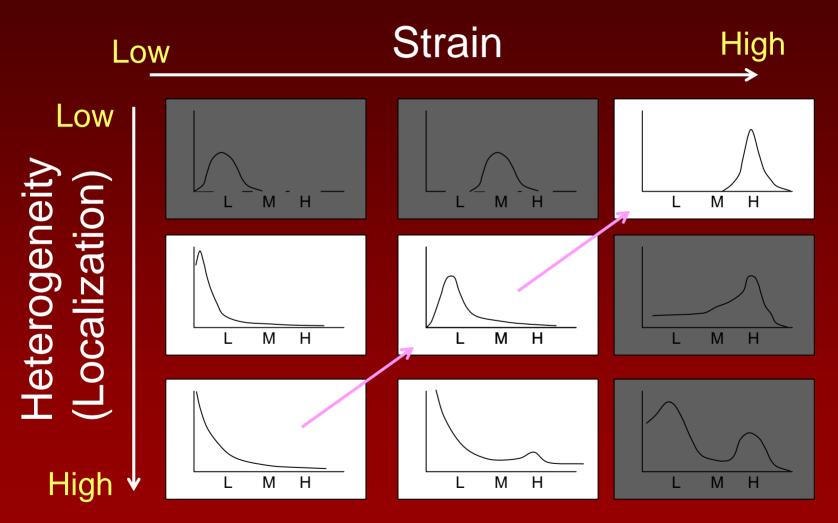
Frequency

Conceptual Variation of Fracture Density Distributions



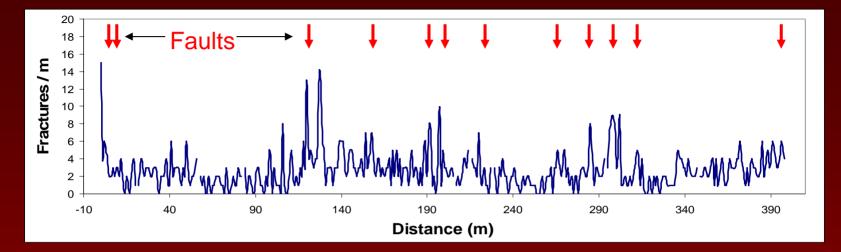


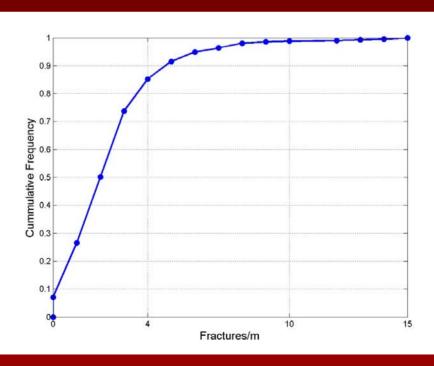
Conceptual Variation of Fracture Density Distributions





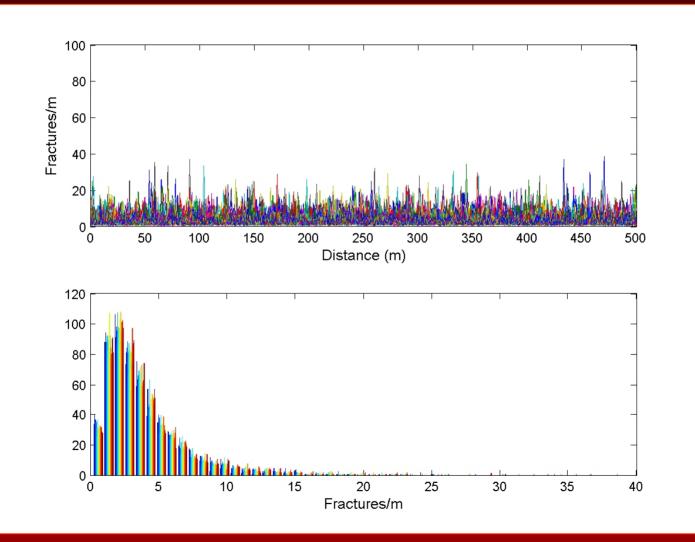
Empirical CDF for Simulation





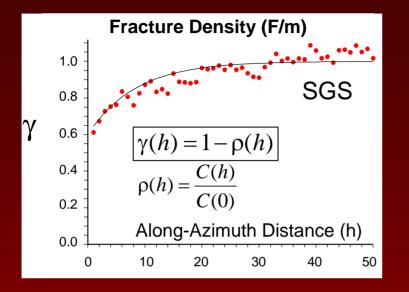


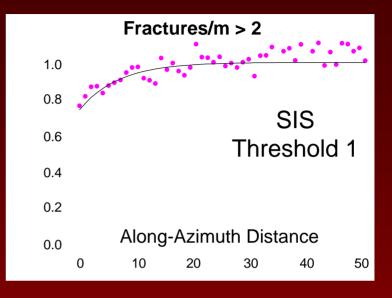
Equally-Probable Realizations

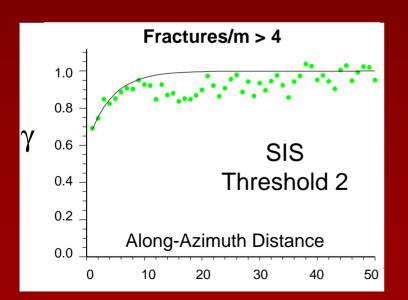


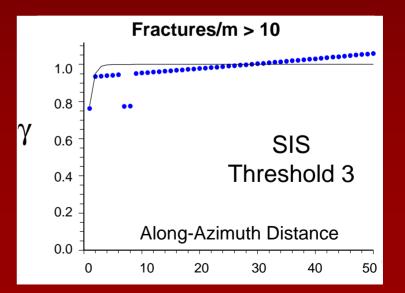


Spatial Correlation of Fracture Frequency





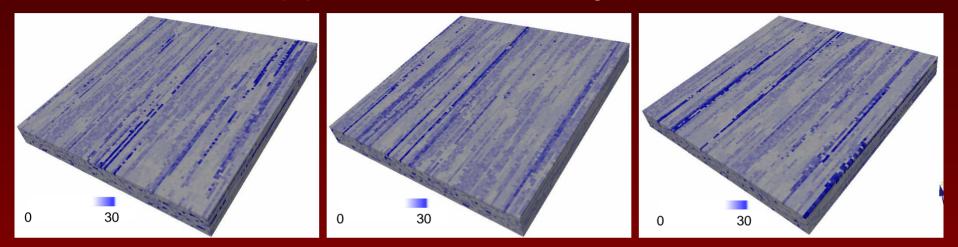




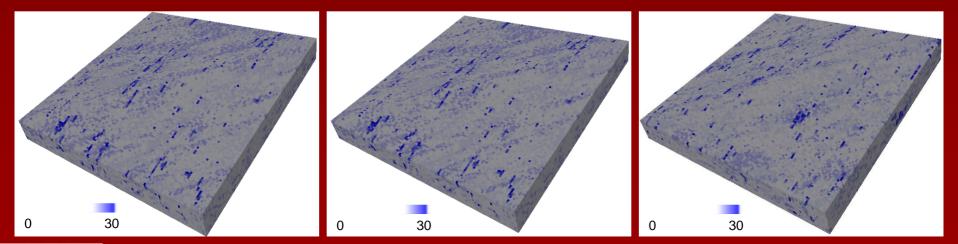


Variogram-Based Sequential Indicator

Dip-parallel Faults and Background Joints



Dip-parallel Faults and Oblique Background Joints (Splays)

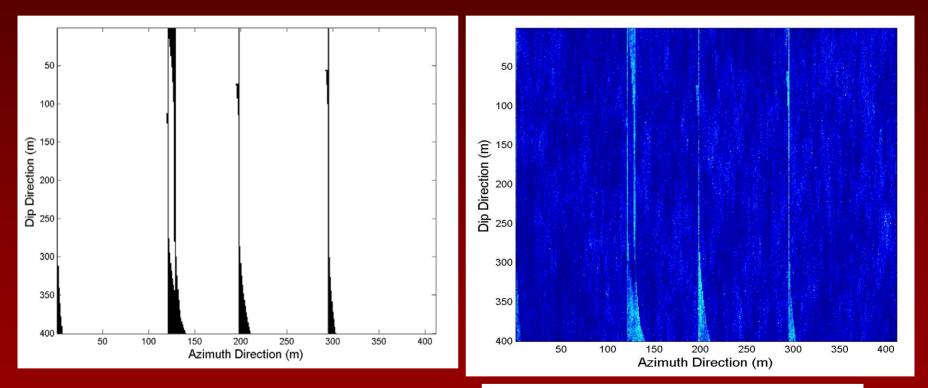




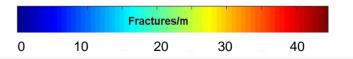
Stochastic Fault Modeling (SFM) and SGS: Object-Based Indicator

Low Shear-Strain Area

Dip-parallel faults and parallel joints



Stochastic Modeling of Fault Distribution and Architecture

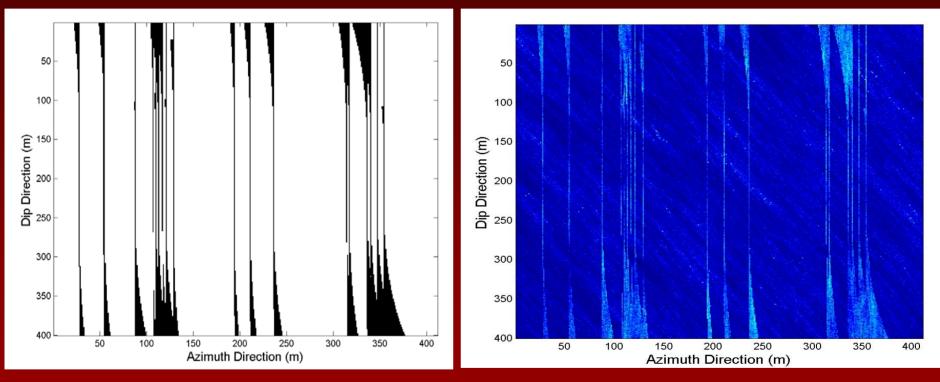


Combination of SFM and SGS: Cookie-cutter Technique



Object-Based Sequential Indicator Simulation (SIS)

Intermediate Shear-Strain Area Dip-parallel faults and oblique splay fractures



Stochastic Modeling of Fault Distribution and Architecture

| Fractures/m | | | | | | |
|-------------|----|----|----|----|----|--|
| 0 | 10 | 20 | 30 | 40 | 50 | |

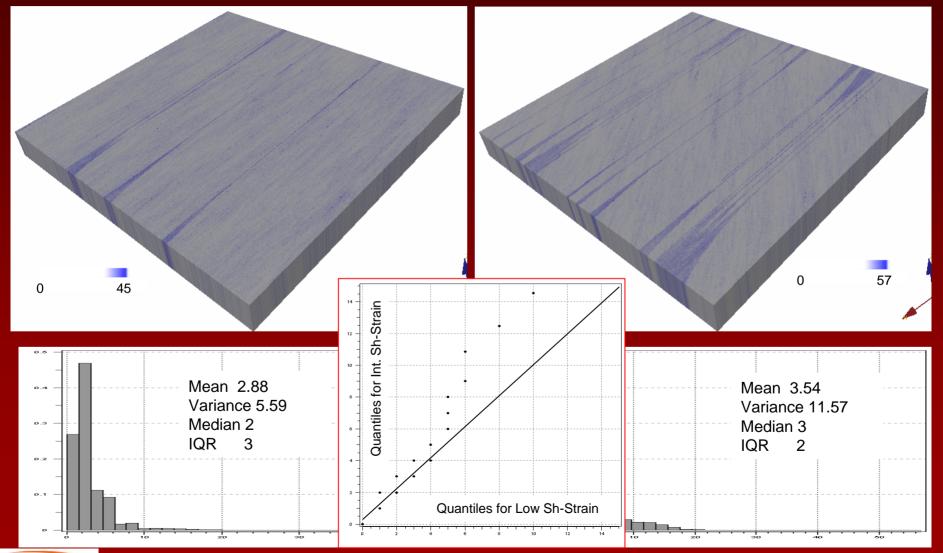
Combination of SFM and SGS: Cookie-cutter Technique



Results from Object-Based Indicator Simulation

Low Shear-Strain Area

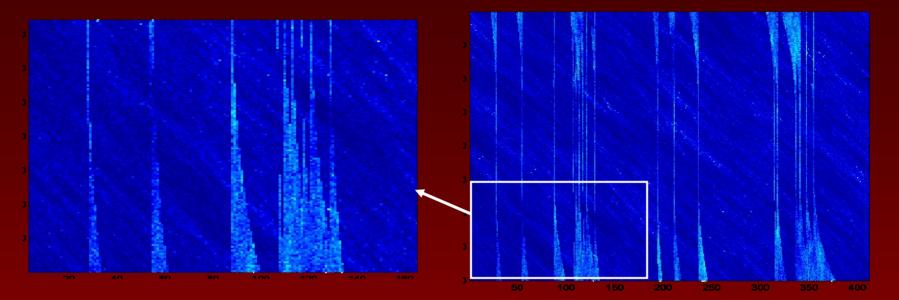
Intermediate Shear-Strain Area





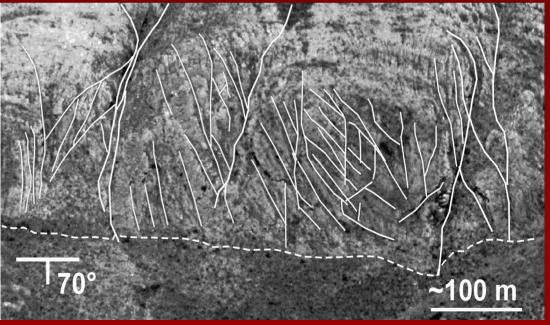
Florez, 2006

Reproducing the Spatial Heterogeneity



Intermediate Shear-Strain Area:

Dip-parallel faults, and oblique splay fractures





Conclusions

- 1. Outcrops provide key information about orientation, deformation mechanisms and spatial heterogeneity of fault and fracture systems.
- 2. In fold and thrust belts fracture systems evolve as the result of hierarchical shearing and progressive deformation, creating different fracture hierarchies.
- 3. There is a 1st order relationship between fracture spacing and stratigraphic architecture.
- 4. Shear strain and lithology are key factors controlling fracture density.
- 5. Antithetic conjugate faults can produce fracture swarms.
- 6. Object-based SIS can be used to model NFR.

